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GROWING SEASON FOR MOISTURE AND TEMPERATURE AVAILABILITY, EVALUATED WITH THE THORNTHWAITE WATER BALANCE IN JALISCO, MEXICO

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All content in this magazine is licensed under a Creative Commons Attribution License. Attribution-Non-Commercial-Non-Derivatives 4.0 International (CC BY-NC-ND 4.0). Abstract: In Mexico, corn is the most important crop produced under rainfed conditions. In this condition, the planning of agricultural operations is highly dependent on rainfall, with lower yields and high risk of crop loss. In recent years, Mexico has been facing increasingly severe, frequent and intense droughts, with important impacts on all economic sectors, but with greater influence on agriculture and livestock. This is the reason that motivates the evaluation of water availability in the growing season (GS) of rainfed crops and to have a quantitative basis to propose management technologies to cushion its effects. The objective of the present work was to evaluate the GS due to the availability of adequate moisture and temperature for crops by means of the Thornthwaite model for Jalisco, Mexico. The study used the water balance model of Thornthwaite and Mather (1957). The characteristics of GS were obtained: onset, termination and duration, according to FAO criteria. The GS characteristics were mapped using interpolation procedure in IDRISI Selva and generating the final presentation in ARCVIEW GIS. The results show that the GS begins before June 10 in the central and western zones of the state; it continues with delays in the north of the state and the highlands of Jalisco until after June 30. The termination of the GS was irregular; it ended in the north and highlands of Jalisco before October 15; it continued in the rest of Jalisco until October 31. The duration of GS was shortest in the extreme north and highlands with less than 115 days, up to less than 160 days in the coast and south of Jalisco.

Keywords: Water balance, rainfed, Jalisco.

INTRODUCTION

In Mexico during 2023, the agricultural area planted with irrigation was 1.906 million hectares and only 1.865 million hectares were harvested, representing a loss of 41 thousand hectares; in rainfed agriculture, 7.395 million hectares were planted and only 6.278 million hectares were harvested, indicating losses of 1.117 million hectares (SIAP-SADER, 2024). The above implies that, in rainy season, in addition to having lower yields, the risk of crop loss is great. According to INEGI (2018), the main cause of crop loss is due to the effect of climatic phenomena not controlled by man, such as droughts, floods, frost and hailstorms. In recent years, Mexico has been facing increasingly severe, frequent and intense droughts, with significant impacts on all economic sectors, but with greater influence on agriculture and livestock (Banco de México, 2022). This is the reason why the availability of water for rainfed crops must be evaluated and have a quantitative basis to propose management technologies that cushion its effects (Bradford et al., 2017; Shaw, 1976).

Moisture availability for rainfed crops can be expressed in terms of deficiency, excess or optimum water conditions for plants (Wilhite and Glantz, 1985). Wang (1963) pointed out that drought or excess moisture (or even flooding) is caused by hydrological imbalances involving regional characteristics such as drainage, runoff, topography, soil structure, animal and plant species of the area, as well as meteorological factors. He also indicated that, due to the impossibility of having complete control of drought or excess moisture, protective and preventive measures must be taken to minimize damage to agricultural production. One action is the probabilistic evaluation of the occurrence of these meteorological phenomena.

Under rainfed conditions, crop growth and development is modified when moisture deficiencies occur according to the degree of intensity of the phenomenon. In such a situation, a more accelerated crop development is observed and the accumulation of physiological time is altered. For this reason, Doorenbos and Kassam (1986) affirm that the lack of attention to the satisfaction of the water needs of a crop, expressed by water deficits, can cause negative effects on growth and yield.

The degree to which plant water requirements are met can be assessed by a soil water balance (Chang, 1968; Shaw, 1976; Baier and Robertson, 1966; Smith *et al.*, 1985). The soil water balance can be expressed by the equation: $PP + Irr = ET \pm \Delta HS + PER + RUN$, where PP is rainfall, Irr is crop irrigation, ET is crop evapotranspiration, ΔHS is soil moisture changes in the root zone, PER is water percolation out of the root zone and *RUN* is water runoff over the soil (Chang, 1968; Smith *et al.*, 1985; Reddy, 1983; Allen et al., 1991).

Quantifying the water required for the growth and development processes of rainfed agricultural crops is a critical part of agronomic management in the field. One of the methods available to estimate the edaphic water balance is that of Thornthwaite and Mather (Thornthwaite and Mather, 1957). Although this method is one of the simplest and most robust, it is feasible to apply it with monthly or decadal or daily data (Thornthwaite and Mather, 1957), it only requires temperature information to calculate the ETP. This method offers the possibility of estimating the water balance at basin (Mammoliti et al., 2021; Moghaddam and Mohammadkhan, 2017), regional (Umar and Yusuf, 2019), global (Mintz and Serafini, 1992; Willmott et al., 1985) scales for any point on the earth's surface (Cardoso do Vale et al., 2022; Umar and Yusuf, 2019), but can also be applied for climate classifications, to define the hydrology of an area and the planning of water use by

crops (Fathurrahman et al., 2022; Fernández-Long et al.,2012; Calvo, 1986).

The use of ETP proposed by Thornthwaite (1948) is based on temperature, with good performance in continental and tropical climates of North America where it was developed; however, its application is limited in arid and semi-arid conditions (Ward and Robinson, 2000). Another approach for calculating ETP is with the use of evaporation measured in the evaporimeter tank type "A", which allows measuring the effects of solar radiation, wind, temperature and air humidity integrated in the evaporation measurement of a free water surface, which analogous to the plant responds to the same climatic factors (Allen et al., 2006; Doorenbos and Pruit, 1977). For these reasons, to improve the performance of the Thornthwaite water balance, the calculation of ETP by evaporation has been proposed (Sabziparvar et al., 2010; Steenhuis and Van der Molen, 1986). To give spatial expression to point characteristics of the water balance, interpolation methods provide accurate estimation at intermediate points (Hartkamp et al., 1999); one of the most widely used is Inverse Distance Weighting (IDW) (Li and Heap, 2011). The objective of the present work was to evaluate the GS for availability of adequate moisture and temperature for crops by means of the Thornthwaite model for Jalisco, Mexico.

MATERIALS AND METHODS

Monthly normal evaporation and precipitation data from 130 climatological stations in the state of Jalisco were used, with distribution shown in Figure 1. These data were obtained from the National Water Commission and published as climatological normals by Flores et al. (2022). This publication also includes the date of first frost, information that was used to correct the duration of the GS.



Geographical location of the 130 climatological stations in the State of Jalisco used in this study.

In the present study, the water balance model of Thornthwaite and Mather (1957) was used, expressed by the following equation: P - ETP $\pm \Delta RH + DEF$ - EXC = 0, where P is precipitation, ETP is potential evapotranspiration, ΔRH is the change in soil moisture reserve. DEF is the water deficit and EXC is excess water. ETP was calculated based on monthly evaporation (EVAP) corrected by the vat coefficient of 0.8, according to the expression: ETP=kp*EVAP. Actual evapotranspiration (ETR) was consumed by the crop. PPutil is equal to (PP - ETP) and indicates the amount of ETP consumed. Soil moisture reserves (RES) occur if the soil storage capacity defined by its physical characteristics (SC) and PP + ETP is not exceeded. Excess water (EXC) was defined when PP + RH exceeded ETP + SC. Water deficit (DEF) occurred when ETP was greater than PP + RH. This procedure was implemented in the EXCEL spreadsheet, as shown in Table 1.

The characteristics of CD: onset, termination and duration were determined according to the FAO definition, expressed by the following conditions: 1) the onset of GS when PP > 0.5 ETP, 2) the termination of GS when PP + RH < 0.5 ETP, adjusted with the date of first frost (1FRO), so that if the 1FRO occurred when moisture was still present it marked the termination of the GS and 3) the duration of the GS is the difference between the termination and the beginning of the GS (Arteaga-Ramírez et al., 2006). The GS characteristics were mapped using IDW interpolation procedure in IDRISI Selva and generating the final presentation in ARCVIEW GIS.

RESULTS AND DISCUSSION

Figure 2 shows the onset of GS (GSons) for Jalisco. This figure shows that the ECini begins before June 10 in the central and western zones of Jalisco, and continues from June 10 to 20 in the coastal, southern, central, southern and middle zones of the highlands region of Jalisco. The GSons is delayed in the north and highlands of Jalisco until after June 30. These results show a trend similar to those reported by Ruíz et al. (2000), who analyzed the characteristics of the GS for Jalisco in two periods: from 1947 to 1971 and 1972 to 1997, although differences can be attributed to the historical series used, the normal data used and the method for defining GSons.



Figure 2. Beginning of the growing season by moisture availability and temperature for Jalisco.

MES	JAN	FEB	MAR	APR	MAY	JUN	JUL	AGO	SEP	OCT	NOV	DEC
EVAP	125.4	149	209.9	239.6	255.9	189.3	156.7	141.1	121.1	117.7	113.2	104.9
PP	17.6	14.6	6.2	5.5	34	174.6	227.7	194	147.9	60.5	13.6	11.5
ETP	100.3	119.2	167.9	191.7	204.7	151.4	125.4	112.9	96.9	94.2	90.6	83.9
ETR	17.6	14.6	6.2	5.5	34	151.4	125.4	112.9	96.9	94.2	51.1	11.6
Useful	-82.7	-104.6	-161.7	-186.2	-170.7	23.2	102.3	81.1	51.0	-33.7	-77.0	-72.4
RES	0	0	0	0	0	23.2	71.2	71.2	71.2	37.5	0	0
DEF	82.7	104.6	161.7	186.2	170.7	0	0	0	0	0	39.5	72.3
EXC	0	0	0	0	0	0	0	54.3	81.1	51	0	0

Table 1. Parameters used in the Thornthwaite water balance.

Figure 3 shows the termination of GS (GSter) in Jalisco. The GSter is more irregular than the beginning of the GS; the end of the GS begins in northern and northeastern areas of Jalisco before October 15; it continues in the rest of northern, highlands and southern areas of Jalisco until October 31. For the Central, South and Ciénega regions of Jalisco, the GSter ends until November 15. The GSter is delayed until November 30 in the Coast and Southern regions of Jalisco, while some areas of the Coast and Southern Jalisco occur after December 1. Comparison of these results with those reported by Ruíz et al. (2000) shows a similar general trend, but with later dates. These differences are also attributed to the historical series used, the use of normal data and the method for defining GSter.

The duration of the GS (GSdur) in Jalisco is shown in Figure 4. This characteristic is defined by the difference between the beginning and the end of the GS. The GSdur is shorter in the extreme north of the state and some zones of the highlands, with less than 115 days; in the middle part of the north and highlands of Jalisco it ranges from 115 to 130 days; for the southern part of the northeastern zone and the southern part of the northern zone of Jalisco, as well as some zones of the center the GSdur ranges from 130 to 145 days. In the Central, Ciénega, South and Valles zones of Jalisco, the GSdur ranges from 145 to 160 days, while in the coast and south of Jalisco it reaches more than 160 days. These results contrast with those reported by Ruíz et al. (2000), with a wider current GSter.



Figure 2. Termination of the growing season by moisture and temperature availability for Jalisco.



Figure 4. Length of growing season by moisture availability and temperature for Jalisco.

Studies conducted on the characterization of GS show important trends, particularly with the current climate change to reduce the amount of precipitation, but to increase extreme intensity events (Douville et al., 2021). The duration of GS shows tendency to reduce with late onset of GS, but the termination of GS with low correlation with climatic parameters (Ruíz et al., 2016).

The characteristics of GS studied in the present research represent an important complement to the technological information available for agriculture in Jalisco, as it allows identifying the cycle of crop genotypes feasible for production in the different agricultural zones and the average sowing date (INIFAP, 2017).

CONCLUSIONS

The characteristics of the GS (onset, termination and duration) for Jalisco were identified as an important complement to the technological information available for agriculture in Jalisco. Although the results obtained are associated with climate trends reported in other studies, differences arise due to the effect of climate change, the length of the historical series of data used, the use of normal data and the parameters used to estimate the GS.

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